

**CASCADING LAKES, SPILLWAYS, AND SUBAQUEOUS DEBRIS FLOWS: MODELS FOR MARTIAN PALEOLAKES.** J.W. Rice, Jr.<sup>1</sup>, V.R. Baker<sup>1</sup>, and D.H. Scott<sup>2</sup>, <sup>1</sup>University of Arizona (Lunar and Planetary Laboratory, 1629 E. University, Tucson AZ 85721, jrjce@lpl.arizona.edu), <sup>2</sup>U.S.G.S. (2255 N. Gemini Dr., Flagstaff, AZ 86001).

*“Certainly, in the future as increased aerial photographic coverage of our celestial neighbors becomes available, the search for, and recognition and study of, ancient lakes, their features and basins, will be of great importance to our space effort in interpretation of past or present celestial environments and in selection of landing sites.”*

C.C. Reeves, Jr., 1968,  
Introduction to Paleolimnology

## **Introduction**

Newly acquired MOLA profiles [1,2], over the northern lowland plains region of Mars, indicate that the regional flatness ( $< 1^\circ$  slopes) and low surface roughness ( $< 2\text{m}$  over  $100\text{ km}$  baseline rms variation) is consistent with large-scale terrestrial depositional environments such as fluvio-lacustrine basins and the oceanic abyssal plains. However, the origin of the flatness of the northern hemisphere is unknown. Current explanations to account for both the extraordinary flatness and smoothness of the northern plains calls for the formation of vast lakes [3,4,5,6,7,8] or oceans [9,10,11]. Recent MOLA results [12] supply additional evidence to corroborate these earlier geomorphic studies which indicated that large standing bodies of water occupied the northern plains of the planet.

## **Cascading Lakes and Spillway Systems**

### Earth:

The recent deglaciation of Eurasia and North America provide us with “Martian-Scale” analogs to extensive catastrophic flood regimes and outwash plains; lacustrine basins, and inter-basin spillways. These large glacial lakes which received flood discharges have huge channels at their inlets and outlets but lack such channels across their bottoms [13]. Additionally, the arrival of floodwaters from topographically higher lakes to downstream lakes triggers a domino-like sequence of lake drainage floods [14]. The enormous volume of flood water exceeded the capacity of the basins, which caused spillover and down-cutting at outlet areas, until the upstream lakes drained. The end result would be massive ponding in the terminal, lowest most topographic basin(s).

The largest cascading lake and spillway system is the Trans-Siberian drainage complex. This system extended from the Chersky Mts. of Siberia in the east to the Alps in the west, and drained a catchment area equaling  $2.3 \times 10^7 \text{ km}^2$  [15]. This is an area 3 times the size of the Amazon basin. This system is comprised of the Black,

Caspian, and Aral Seas as well as several lakes. The total areal coverage of these bodies of water exceeded  $3 \times 10^6 \text{ km}^2$ . This flood waters focused first in the Black Sea and then marched west via the Sea of Marmosa to finally ended up in the Mediterranean Sea and Atlantic Ocean [15].

### Mars:

Scott *et al.*, [3] identified 15 large topographic depressions, each having closures exceeding  $100,000 \text{ km}^2$  in areal extent and depths of  $1000\text{ m}$  or more; these basins were considered prospective sites for paleolakes. Elysium, Amazonis, Utopia, Isidis, and Chryse along with their possible interconnections or spillways are the five basins that contain the best evidence of standing bodies of water. These basins drain an estimated total area of  $4.8 \times 10^7 \text{ km}^2$ . The total areal coverage of these bodies of water is estimated to be  $7 \times 10^6 \text{ km}^2$ . Individual lake basins on Mars (Chryse, Utopia, Isidis, Amazonis, and Elysium) were either / both the precursors to the northern polar ocean or last vestiges of the ocean as it dried up. The cascading lake scenario could have initially provided the flood water impetus to form the northern ocean.

The Chryse basin clearly exhibits huge inlet channels, outlet channels and lacks any incised channels across its floor. Maja Valles flows in from the southwest, Kasei Valles from northwest, Shalbatana and Simud Vallis from south, Tiu and Ares Vallis from the southeast, and Mawrth Vallis from the northeast. These morphologic associations are observed in large terrestrial lake basins which receive enormous flood discharges and eventually cascade or spill over into another basin [13]. The Chryse basin floor contains an area of undissected terrain with wrinkle ridges ( $2.75 \times 10^3 \text{ km}^2$ ). These wrinkle ridges appear subdued and mantled in places perhaps due to ponding and burial by flood sediments. Indications of ponding and subsequent outflow are also observed in the northeast of Chryse Planitia ( $30^\circ\text{N}$ ;  $41^\circ\text{W}$ ), where streamlined islands are ori-

ented to flow from Maja Valles, some 950 km away.

### Subaqueous debris flows

#### Earth:

Daly [16] was the first to suggest the idea of turbidity currents as a geologic agent. He proposed that these sediment laden currents were responsible for the origin of submarine canyons. The first paper to accurately describe an actual subaqueous debris flow, turbidity current and submarine slump was based on the 1929 Grand Banks earthquake [17]. This earthquake shook the continental slope south of Newfoundland and resulted in the destruction of all the submarine telegraph cables lying downslope of the epicentral area. Heezen and Ewing [17] explained the successive downslope cable obliteration by means of a subaqueous debris flow which transitioned into a turbidity current. This sediment flow then rumbled down the continental slope and eventually settled out onto the abyssal plains, 750 km from the continental slope. Heezen, et al. [18] determined that the flat floors of the oceanic abyssal plains are due to the ponding of sediments derived from turbidity currents. The oceanic abyssal plains are the flattest surfaces on Earth.

Subaqueous debris flows can travel over slopes as low as  $0.1^\circ$  and a distance of 1,600 km [19]. In the Canary Basin, subaqueous debris flow deposits cover an area of 30,000 km<sup>2</sup>. Similar deposits have also been observed on the Amazon cone, North American continental rise, Gulf of Mexico, and Mediterranean Sea. Subaqueous debris flows are a much more important depositional process on the sea floor than previously suspected. For example, subaqueous debris flows generally develop into turbidity currents.

#### Mars:

Recent work by Tanaka [20, 21] proposes that subaerial debris flows formed the Hesperian smooth plains Simud/Tiu deposits. [20, 21] geologically mapped the whole southeast region of Chryse Planitia and studied the regional morphology. Subaerial debris flows were the primary mechanism that Tanaka [20, 21] investigated, he gave the subaqueous debris flow hypothesis only a cursory mention.

We propose that the deposits located near the mouths of outflow channels in the northern plains are subaqueous debris flows that were emplaced in standing bodies of water that collected when the outflow channels debouched into these basins. We will discuss the morphology of terres-

trial subaqueous debris flows and compare this to what we observe on Mars.

### References

- Reeves, C.C., Jr., *Introduction to Paleolimnology*, Elsevier Publ. Co., New York, 228p., 1968. [1] Aharonson, O., M.T. Zuber, G.A. Neumann, and J.W. Head III, Mars: Northern hemisphere slopes and slope determinations, *Geophys. Res. Letts.*, 25, No. 24, 4413-4416, 1998. [2] Smith, D.E., M.T. Zuber, H.V. Frey, J.B. Garvin, J.W. Head, D.O. Muhleman, G.H. Pettengill, R.J. Phillips, S.C. Solomon, H.J. Zwally, W.B. Banerdt, and T.C. Duxbury, Topography of the northern hemisphere of Mars from MOLA, *Science*, 279, 1686-1692, 1998. [3] Scott, D. H., J. W. Rice, Jr., and J. M. Dohm, Martian paleolakes and waterways: Exobiologic implications, *Origin of Life and Evolution of the Biosphere*, 21, 189-198, 1991. [4] Scott D.H., M.G. Chapman, J.W. Rice, Jr., and J.M. Dohm, New Evidence of Lacustrine Basins on Mars: Amazonis and Utopia Planitiae, *Proceedings of Lunar and Planetary Science Conference XXII*, 53-62, 1992. [5] Scott, D. H., J. M. Dohm, and J. W. Rice, Jr., Map showing channels and possible paleolake basins, *U.S. Geol Surv. Misc. Inv. Ser. Map I-2461*, 1995. [6] Rice, J.W., Jr., Antarctic Lakes (Above and Beneath the Ice Sheet): Analogs for Mars, in *MSATT Workshop on the Polar Regions of Mars: Geology, Glaciology, and Climate History*, LPI Tech. Rept. 92-08, Part I, 23-24, 1992. [7] Rice, J.W., Jr., Polar Beach Processes, *Lunar Planetary Science Conference XXV*, p.1123-1124, 1994. [8] Rice, J.W., Jr., *Aqueous Sedimentary Basins On Mars*, Ph.D. Dissertation, Arizona State University, 134p., 1997b. [9] Parker, T. J., R. S. Saunders, and D. M. Schneeberger, Transitional morphology in west Deuteronilus Mensae, Mars: Implications for modification of the lowland/upland boundary, *Icarus*, 82, 111-135, 1989. [10] Parker, T. J., D. S. Gorsline, R. S. Saunders, D. C. Pieri, and D. M. Schneeberger, Coastal geomorphology of the martian northern plains, *J. Geophys. Res.*, 98, 11,061-11,078, 1993. [11] Baker, V. R., R. G. Strom, V. C. Gulick, J. S. Kargel, G. Komatsu, and V. S. Kale, Ancient oceans, ice sheets and the hydrological cycle of Mars, *Nature*, 352, 589-594, 1991. [12] Head, III, J.W., M. Kreslavsky, H. Hiesinger, M. Ivanov, S. Pratt, N. Seibert, D.E. Smith, and M.T. Zuber, Oceans in the past history of Mars: Tests for their presence using MOLA data, *Geophys. Res. Letts.*, 25, No. 24, 4401-4404, 1998. [13] Kehew, A.E. and L. Clayton, Late Wisconsin floods and development of the Souris-Pembina spillway system in Saskatchewan, North Dakota, and Manitoba. In: *Glacial Lake Agassiz* (eds: J.T. Teller, and L. Clayton), Geol. Assoc. Canada Spec. Paper, 26, 187-209, 1983. [14] Kehew, A.E and M.L. Lord, Origin and large scale erosional features of glacial lake spillways in the northern Great Plains, *Geol. Soc. Am. Bull.*, 97, 162-177, 1986. [15] Grosswald, M.G., New approach to the Ice Age paleohydrology of northern Eurasia. In: *Palaeohydrology and environmental change* (eds. G. Benito, V.R. Baker, K.J. Gregory), Wiley and Sons, New York, 353p., 1998. [16] Daly, R.A., Origin of submarine canyons, *Am. J. Sci.*, 31, 5<sup>th</sup> Ser., 401-420, 1936. [17] Heezen, B.C. and M. Ewing, Turbidity currents and submarine slumps, and the 1929 Grand Bahamas earthquake, *Am. J. Sci.*, 250, 849-873, 1952. [18] Heezen, B.C., M. Ewing, and D.B. Ericson, Submarine topography in the North Atlantic, *Geol. Soc. Am. Bull.*, 62, 1407-1409, 1951. [19] Embley, R.W., New evidence for the occurrence of debris flows in the deep sea, *Geology*, 4, 371-374, 1976. [20] Tanaka, K.L., Sedimentary history and mass flow structures of Chryse and Acidalia Planitiae, Mars, *J. Geophys. Res.*, 102, 4131-4149, 1997. [21] Tanaka, K.L., Debris-flow origin for the Simud/Tiu deposit on Mars, *J. Geophys. Res.*, 104, No. E4, 8637-8652, 1999.